

Renewable Energy Systems (12210588)

## 6. Solar Radiation

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Lecture 6

Earth Energy Budget

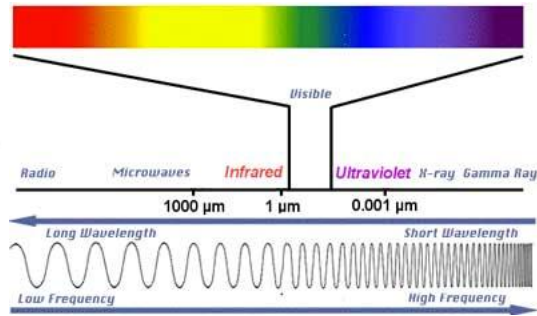
# Net Radiation

Net R = net SW – net LW

$$R_{NET} = \downarrow_{SW} + \downarrow_{SW} - \uparrow_{SW} - \uparrow_{LW} + \downarrow_{LW}$$

where

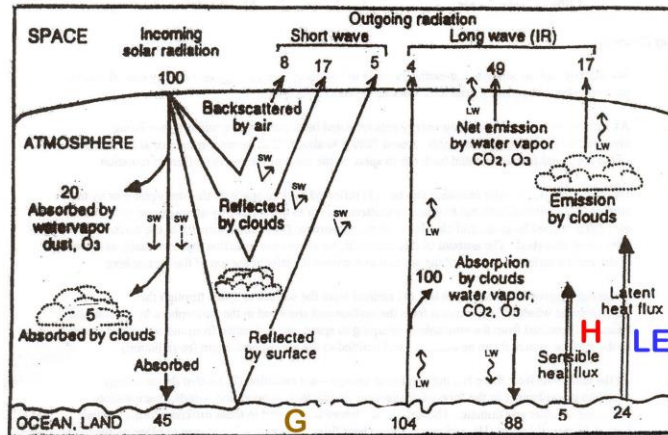
SW = direct SW + diffuse SW



## Shortwave and Longwave Radiation

- Visible light and ultraviolet radiation are commonly called shortwave radiation, while infrared radiation is referred to as longwave radiation.
- The Sun radiates energy mainly in the form of **visible light**, with small amounts of ultraviolet and infrared radiation.
- For this reason, solar radiation is usually considered **shortwave** radiation.

<http://www.kidsgeo.com/geography-for-kids/0060-shortwave-longwave-radiation.php>

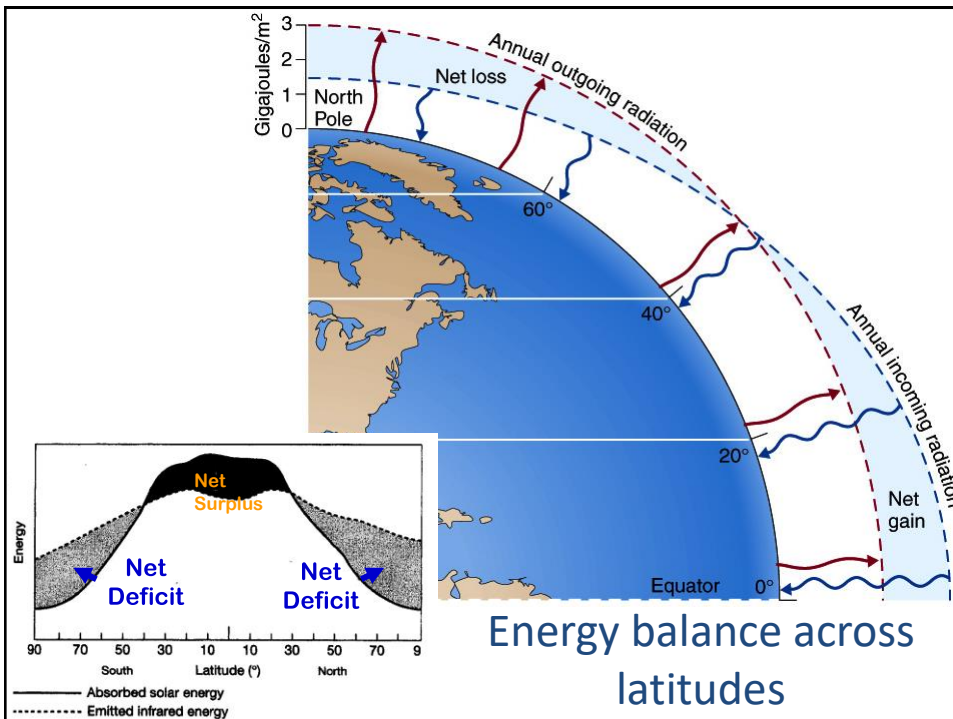
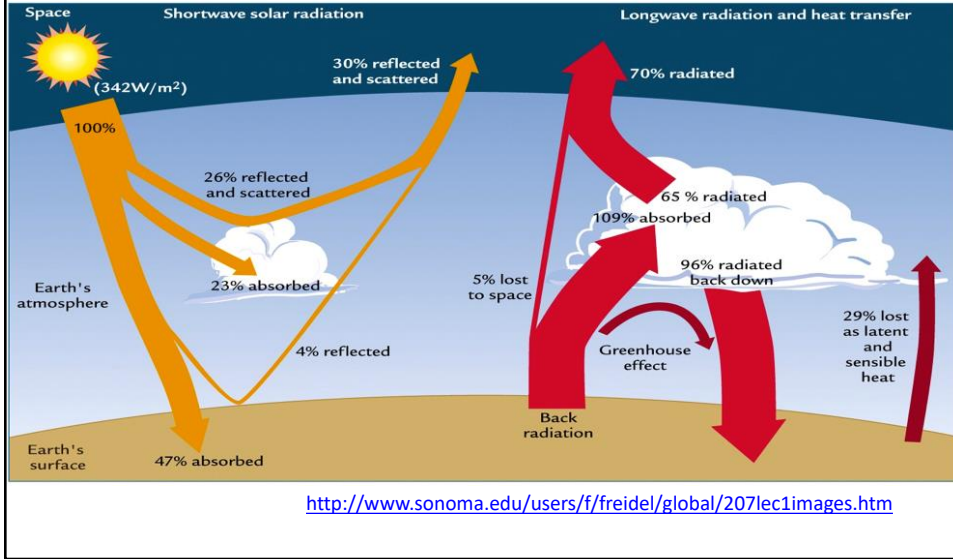


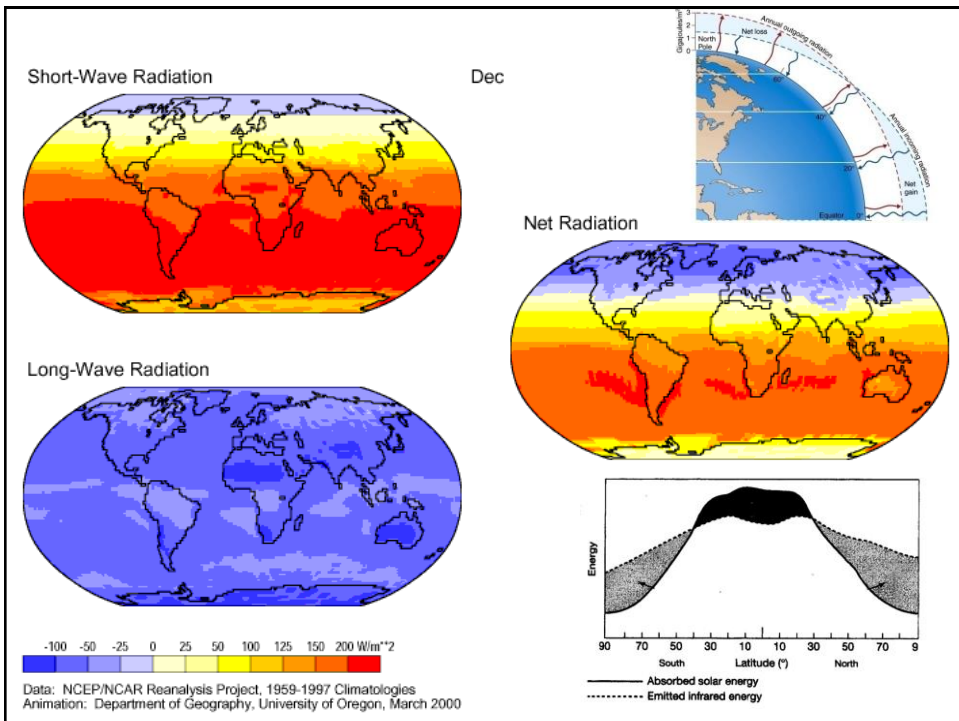
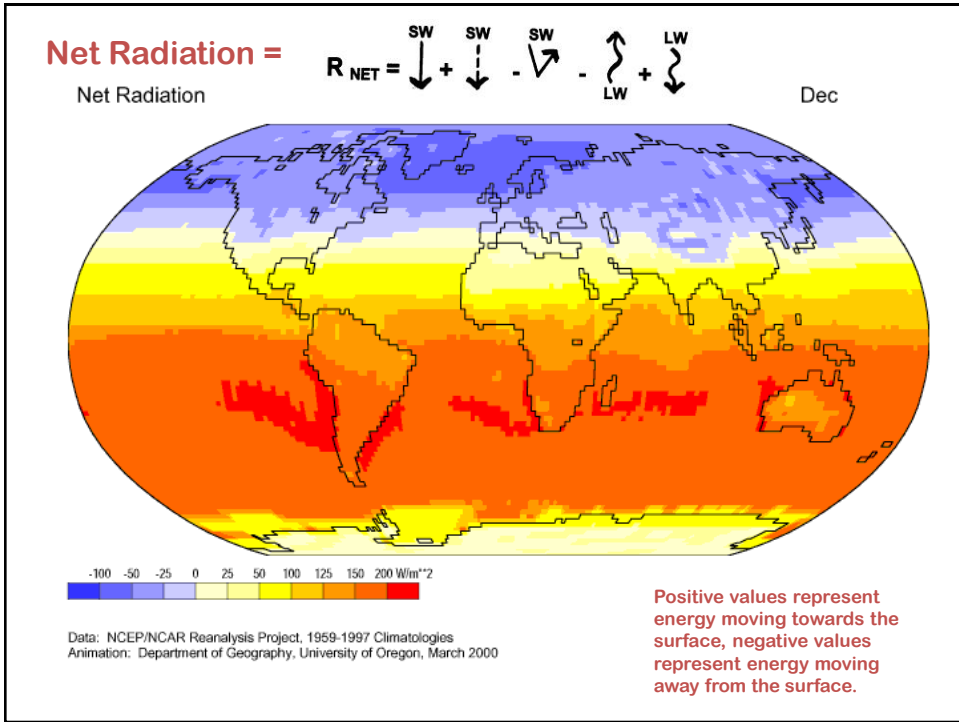
## Solar energy budget

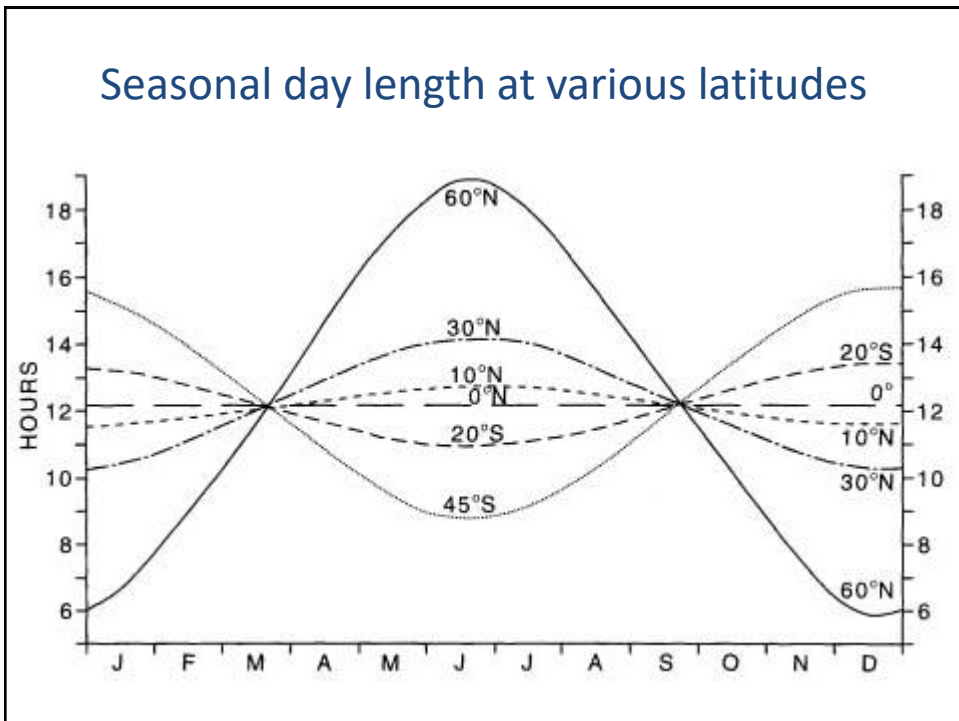
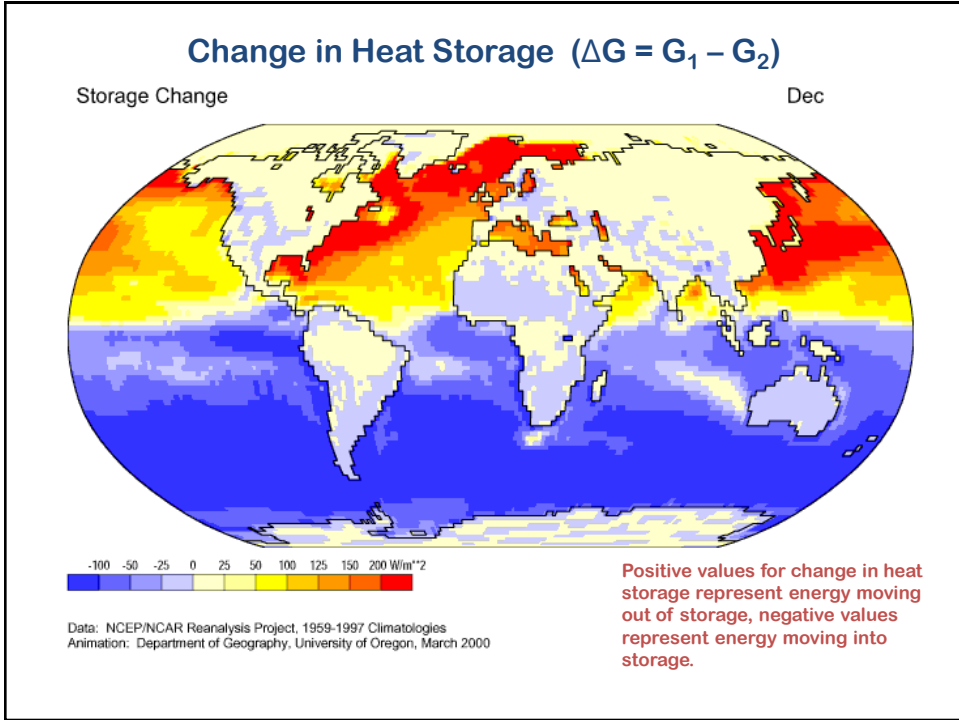
	IN		OUT		R <sub>NET</sub> IN - OUT
	SW	LW	SW	LW	
Balance at top of atmosphere	100	0	8 + 17 + 5 = 30	4 + 49 + 17 = 70	0
Balance within atmosphere	20 + 5 = 25	100	0	88 + 49 + 17 = 154	- 29
Balance at Earth's surface	45	88	0	104	+ 29

R<sub>NET</sub> = 0 at the top of the atmosphere  
 R<sub>NET</sub> = + 29 at the Earth's surface, but these surplus units go into H + LE fluxes to atmosphere (G = 0)  
 R<sub>NET</sub> = - 29 within the atmosphere, but this deficit is balanced by H + LE fluxes from Earth's surface

# Earth's energy balance





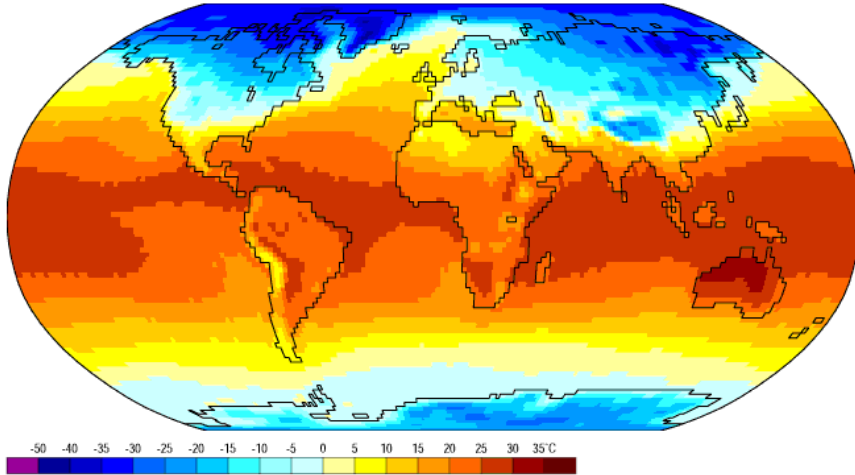


## Air Temperature (at the surface) = T (°C)

Seasonal temperature variations can be explained in terms of the latitudinal & seasonal variations in the surface energy balance.

Air Temperature

Dec



Data: NCEP/NCAR Reanalysis Project, 1959-1997 Climatologies  
Animation: Department of Geography, University of Oregon, March 2000

The pattern of temperatures is a function of net short-wave radiation, net long-wave radiation, sensible heat flux, latent heat flux and change in heat storage.

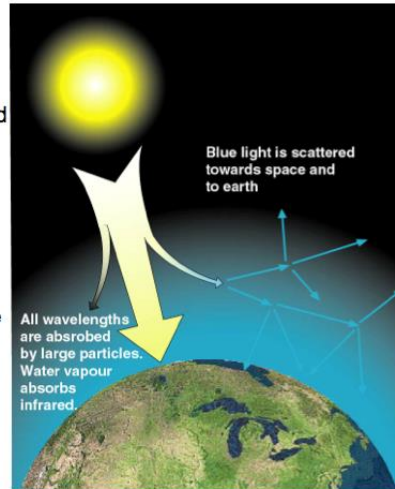
## Solar radiation and position of the sun

## Terrestrial Radiation

Atmosphere has several different effects on radiation:

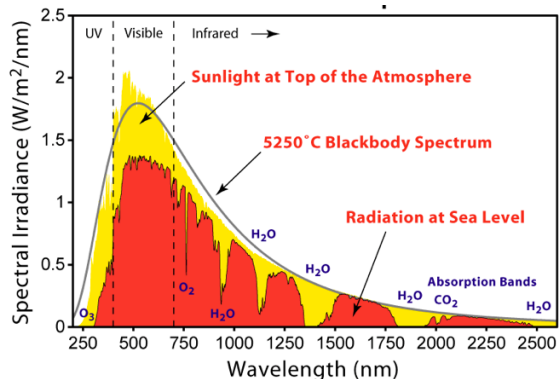
1. scattering, 2. absorption, and 3. reflection.

- Earth's atmosphere has several impacts on radiation:
  - Scattering of ~10% of light causes this light to hit earth's surface at a wide range of angles and coming from anywhere in the sky. It is most effective for higher energy photons.
    - Direct light is the light from the sun which reaches the earth without scattering.
    - Diffuse light is scattered by the atmosphere.
  - Absorption in the atmosphere changes both the power density and the spectral distribution of terrestrial solar spectrum.
    - Ozone absorbs at high photon energies.
    - Water vapor, CO<sub>2</sub>, absorb in infra-red.
  - Clouds, other local variation in atmosphere introduce variability (both locally and temporally) into terrestrial solar radiation.



## Solar radiation spectrum

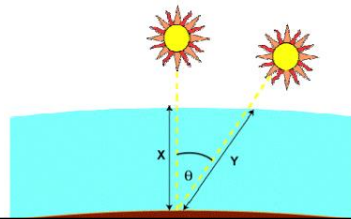
- As the sunlight travels through the atmosphere, chemicals **interact** with the sunlight and **absorb** certain wavelengths
- Example: O<sub>3</sub> strips UV light
- Atmospheric **scattering** plays a role, removing higher frequencies from direct sunlight and scattering it about the sky
- This is why the **sky** appears **blue** and the **sun yellow**
- The greater the distance in the atmosphere through which the sunlight travels, the greater this effect
- This is why the **sun** looks **orange** or **red** at dawn and sundown



[http://en.wikipedia.org/wiki/File:Solar\\_Spectrum.png](http://en.wikipedia.org/wiki/File:Solar_Spectrum.png)

## Air Mass (AM)

- AM defines the direct optical path length through the Earth's atmosphere, expressed as a ratio relative to the path length vertically upwards, i.e. at the zenith
- AM is defined as 
$$AM = \frac{1}{\cos(\theta)}$$
- It is used to characterize the performance of solar cells under standardized conditions, and is often referred to using the syntax "**AM**" followed by a number
- AM1.5**: is almost universal (standard) when characterizing terrestrial power-generating panels
- AM1.5G**: G stands for global, or direct and diffuse radiation
- The solar radiation outside the Earth's atmosphere is called **AM0**

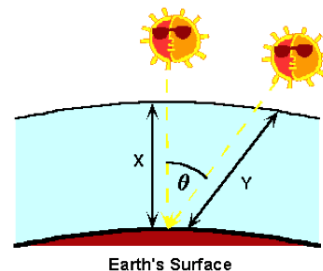


## Direct and global irradiance

- Absorption in incident light is approximated by the concept of Air Mass (**AM**)
- Power density in space is called **AM0**
- Standard conditions **AM1.5**
- Direct and indirect radiative power can be approximated from AM value

$$I_D = 1.353 \times (0.7^{AM})^{0.678}$$

$$I_G = 1.1 \times I_D$$

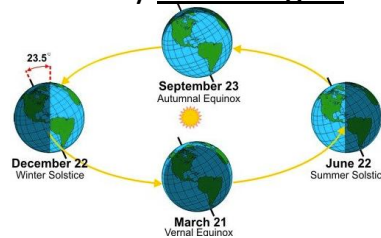


where  $I_D$  is direct irradiance,  $I_G$  is global irradiance (direct + indirect (diffuse))



## Angular dependence of solar radiation

- The **angle** between the **sun** and the **earth's surface** is continually changing because of the rotation of the earth about the sun and the rotation of the earth on its axis
- Determining the **angle** between the sunlight and a surface is critical in finding the **usable sunlight** since only the component of the incident sunlight which is normal to a surface can be used by the surface
- The **position** of the **sun** is specified by three angles:
  - **Declination angle** ( $\delta$ )
  - **Elevation angle** ( $\alpha$ )
  - **Azimuth angle** ( $\theta$ )



<http://www.businessinsider.com/standing-an-egg-on-its-head-has-nothing-to-do-with-the-spring-equinox-2013-3>

## Angular dependence II

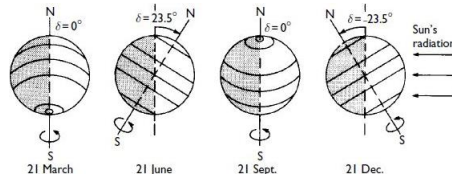
- When calculating solar angles, it is more convenient to express the **day of the year** ( $d$ ) as the number of degrees which the earth has rotated around the sun – denoted by  $\Omega$

$$\text{Daily: } \Omega = \frac{360^\circ}{365}(d - 81) \quad \text{Hourly: } \Omega = \frac{360^\circ}{365} \left( d - 81 + \frac{\text{hour} - 12}{24} \right)$$

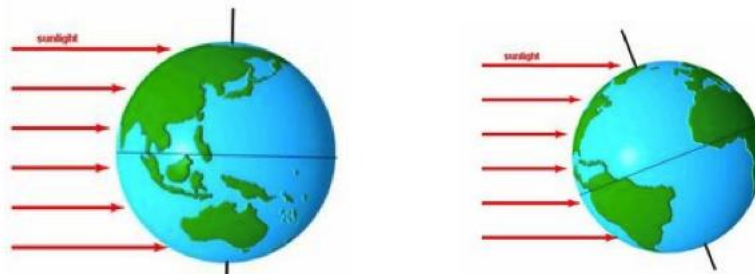
- Angular dependence has a **daily** and **seasonal** component
- **Tilt of the Earth's axis** and **rotation around the sun** causes a variation in the **angle** at which the sun strikes the earth
- **Declination** angle represents this **tilt**
- **Maximum** value of the **declination angle** is **23.45°**

## Rotation of Earth

- Declination angle: angle between the sun and earth

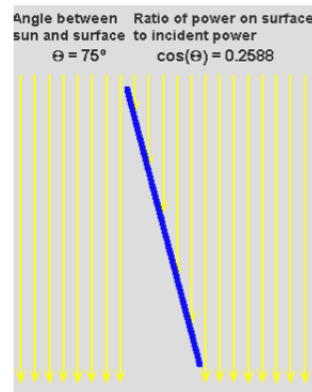
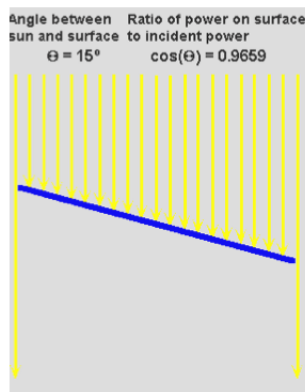


- Declination angle, which is the same everywhere on earth at a given time, changes with seasons



## Earth surface inclination

- Rotation of the earth changes the angle at which the sun strikes the earth surface and though the power density of sunlight on the earth's surface changes



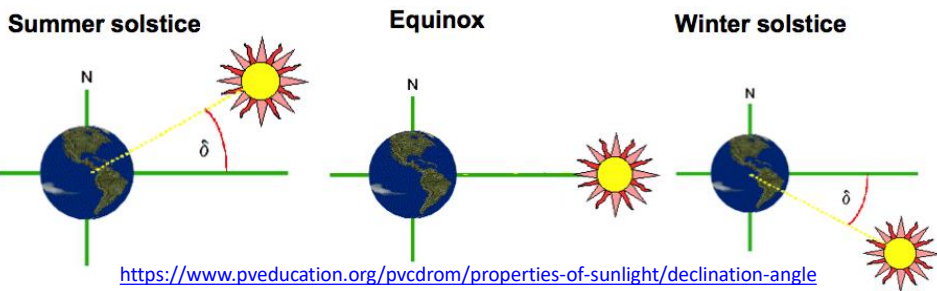
<https://www.pveducation.org/pvcdrom/properties-of-sunlight/declination-angle>

## Declination angle

- **Maximum** and **minimum** declination angle occur at **summer** and **winter** solstice, respectively
- **Declination** angle is **zero** at the **equinoxes**

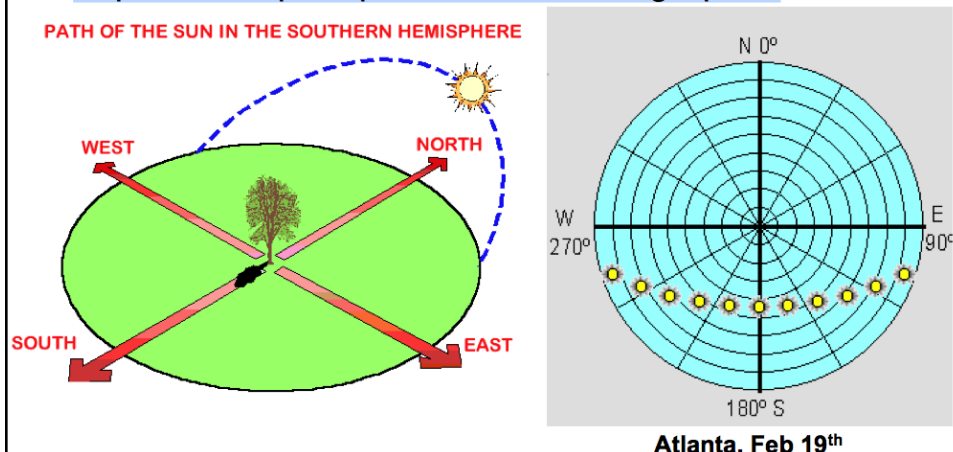
$$\delta = \sin^{-1}[\sin(23.45^\circ)\sin\Omega] \approx 23.45^\circ \sin\Omega$$

$$\delta = 23.45^\circ \sin\left[\frac{360}{365}(284 + d)\right] \quad \text{d: day of the year (DoY)}$$



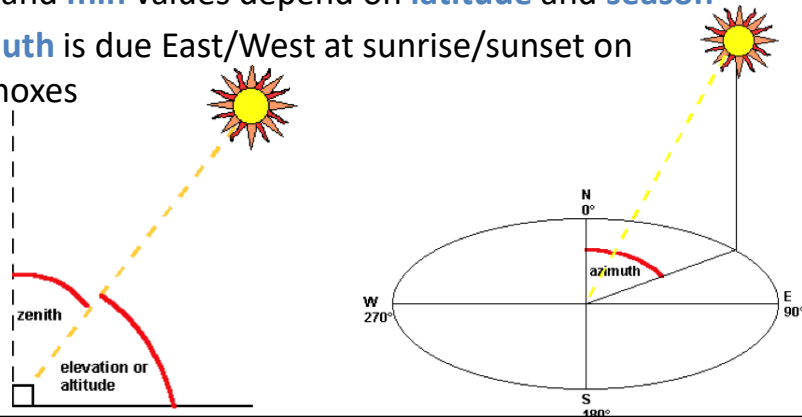
## Path of the sun

- Net **path** of sun throughout the day calculated based on **azimuth** and **altitude** angles
- Represent on **polar plots** or **elevation angle** plots



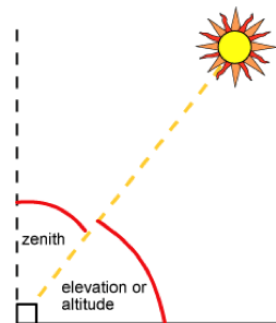
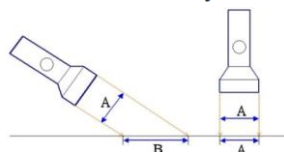
## Azimuth and elevation angles

- Daily and seasonal **angular dependence** at a particular location given by **azimuth** and **elevation** (or zenith or altitude) angles
- Both **elevation** and **azimuth vary** through the day
- **Max** and **min** values depend on **latitude** and **season**
- **Azimuth** is due East/West at sunrise/sunset on equinoxes



## Elevation angle

- The elevation angle (used interchangeably with altitude angle) is the angular height of the sun in the sky measured from the horizontal.
  - Note: both altitude and elevation are also used to describe the height in meters above sea level – this is NOT the same parameter.
- Zenith angle is measured from vertical rather than horizontal.
- The elevation angle varies throughout the day – it is  $0^\circ$  at sunrise and achieves its maximum value at solar noon.
- The elevation angle at solar noon is largest at Summer Solstice (where it equals latitude plus maximum declination ( $23.45^\circ$ )) and lowest at Winter Solstice, (where it equals latitude minus maximum declination ( $23.45^\circ$ )).
- For most many latitudes (greater than  $23^\circ$ ), the sun is never directly overhead.



## Elevation (or altitude) angle

- **Maximum** and **minimum elevation** angle depend on **latitude** and **season**
- **Maximum altitude** angle can be found from **declination** and **latitude**

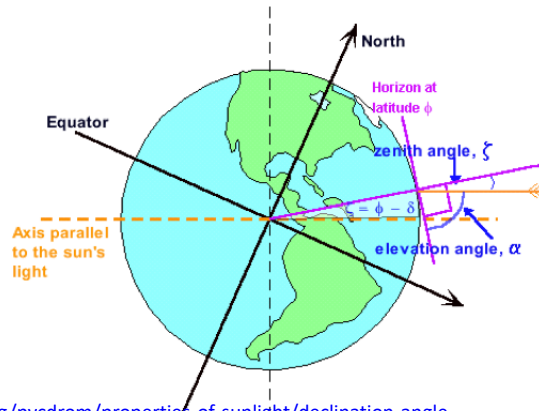
$$\alpha = 90^\circ - (\phi - \delta)$$

where

$\delta$  is the declination angle at a particular day of the year

$\phi$  is the latitude angle

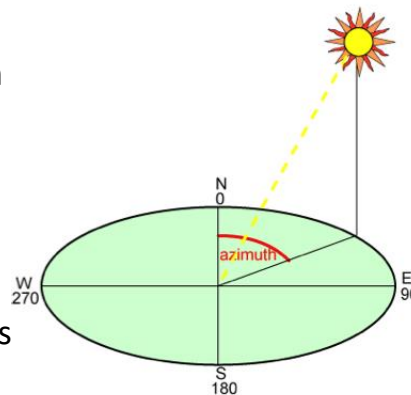
$\alpha$  is the altitude angle



<https://www.pveducation.org/pvc/drom/properties-of-sunlight/declination-angle>

## Azimuth angle

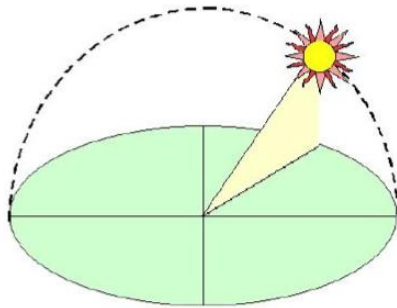
- The **azimuth** angle is the **compass** direction from which the sunlight is coming
- At **solar noon**, the **sun** is always directly **south** in the **northern** hemisphere and directly north in the southern hemisphere
- At the **equinoxes**, the **sun** rises directly **east** and sets directly **west** regardless of the latitude, thus making the **azimuth** angles **90°** at sunrise and **270°** at sunset



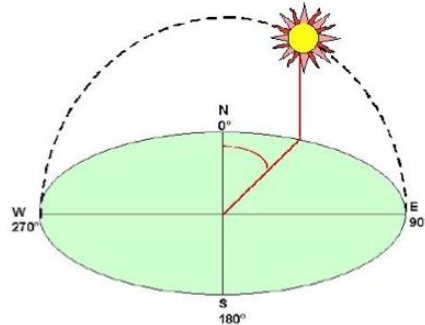
<https://www.pveducation.org/pvc/drom/properties-of-sunlight/declination-angle>

## Solar radiation for hourly calculations

- To calculate the **azimuth** and **elevation** (also called altitude or zenith ( $90 - \text{elevation}$ )) angles throughout each day



Altitude or elevation angle



Azimuth angle

$$\sin \alpha = \sin \phi \sin \delta + \cos \phi \cos \delta \cos(hra)$$

$$\cos \Theta = \frac{\cos \phi \sin \delta - \sin \phi \cos \delta \cos(hra)}{\cos \alpha}$$

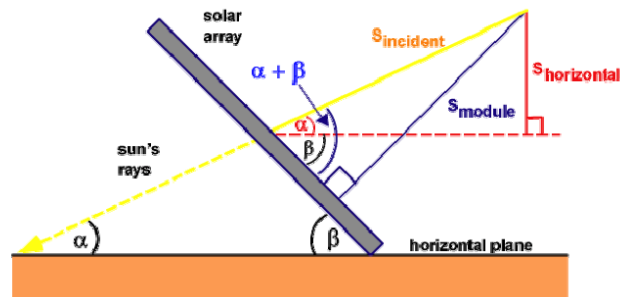
## Power on a tilted surface

- The amount of **solar radiation** incident on a **tilted** module surface is the **component** of incident solar radiation which is **perpendicular** to module surface
- In solar radiation data sets, the **sunlight** is often specified as the component **normal** to a horizontal surface ( $S_{\text{horizontal}}$ )

$$S_{\text{horizontal}} = S_{\text{incident}} \sin(\alpha)$$

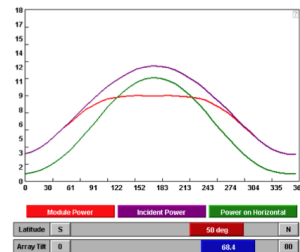
$$S_{\text{module}} = S_{\text{incident}} \sin(\alpha + \beta)$$

$$S_{\text{module}} = \frac{S_{\text{horizontal}} \sin(\alpha + \beta)}{\sin(\alpha)}$$



## Effect of Module Tilt

- **Maximum** power into a module occurs when the **surface** is **normal** to the incident **sunlight**, but since the angle of the sun is continually changing, a module with **fixed tilt** captures only a **fraction** of the total power over a year
- **Optimum** often to set  
 module tilt = latitude  
 but usually get equal amounts of sunlight in summer and winter
- For **maximum power**:  
 module tilt = Latitude - 15°
- For more **uniform production of power**:  
 module tilt = Latitude + 15°



**RENEWABLE  
ENERGY**  
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Israel's Renewable Energy Sector  
at a Crossroads

- Israel transfers 290 MW renewable energy quotas to the solar PV sector as follows:
  - 70 MW of large wind power parks
  - 20 MW of small wind power installations (up to 50 kW each)
  - 200 MW of solar thermal power plants
- **Target:**
  - 10% electricity generation from renewable energy by 2020 and this will lead to savings of more than 2 billion NIS
  - 17 – 30% electricity generation from renewable energy by 2030

[www.renewableenergyworld.com](http://www.renewableenergyworld.com)

<https://www.timesofisrael.com/israel-ups-2030-renewable-energy-target-from-17-to-30-at-cost-of-22-billion/>